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15.023J / 12.848J / ESD.128J Global Climate Change: Economics, Science, and Policy
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The Mathematics of Climate Modeling

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OUTLINE



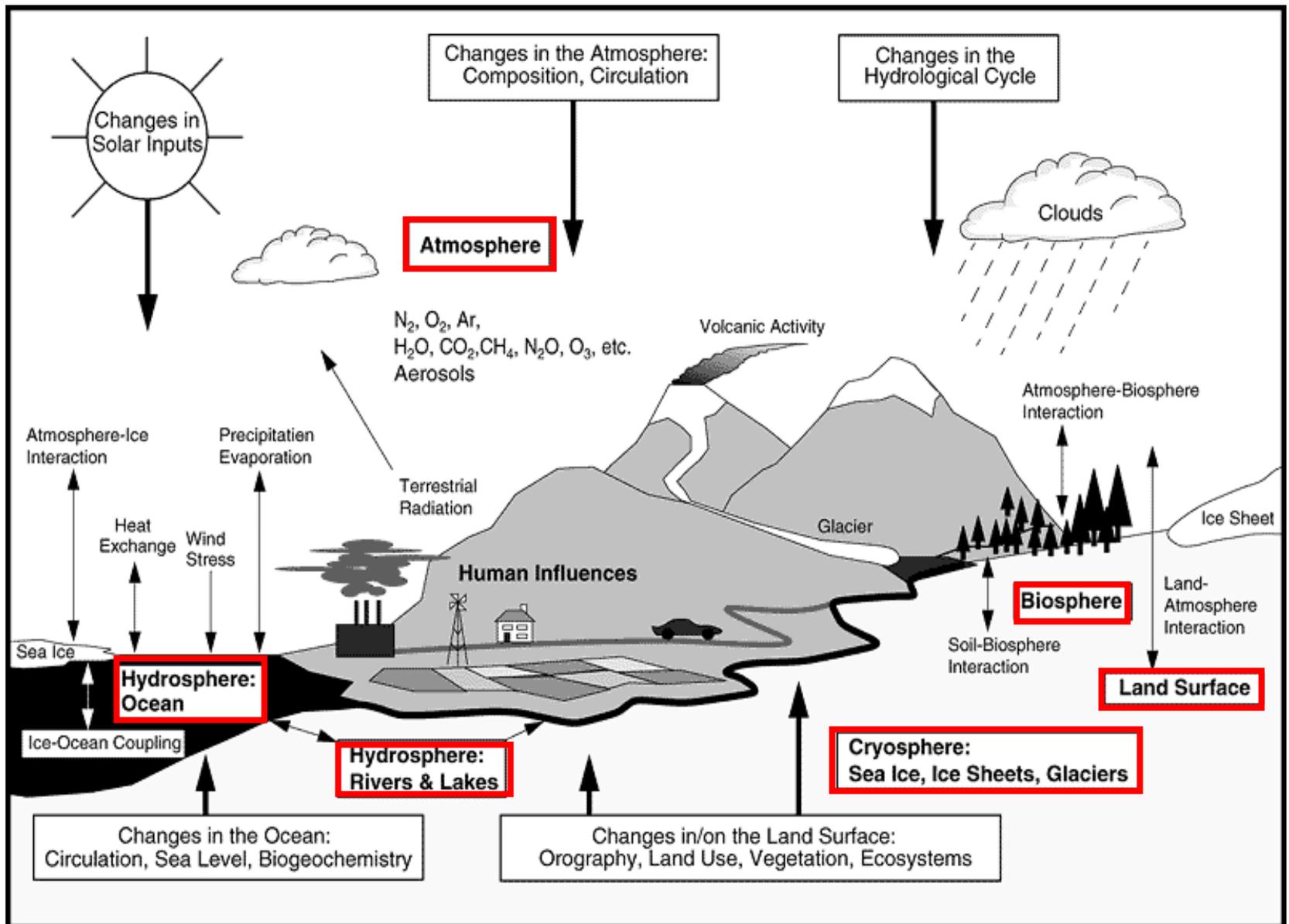
1. What is a **CLIMATE MODEL**?
2. Designing a model
 - Spatial Grid
 - Continuity equation
 - Time step and stability
3. Solving the equations
 - Reality... Computation time and parameterization



1. What is a CLIMATE MODEL?



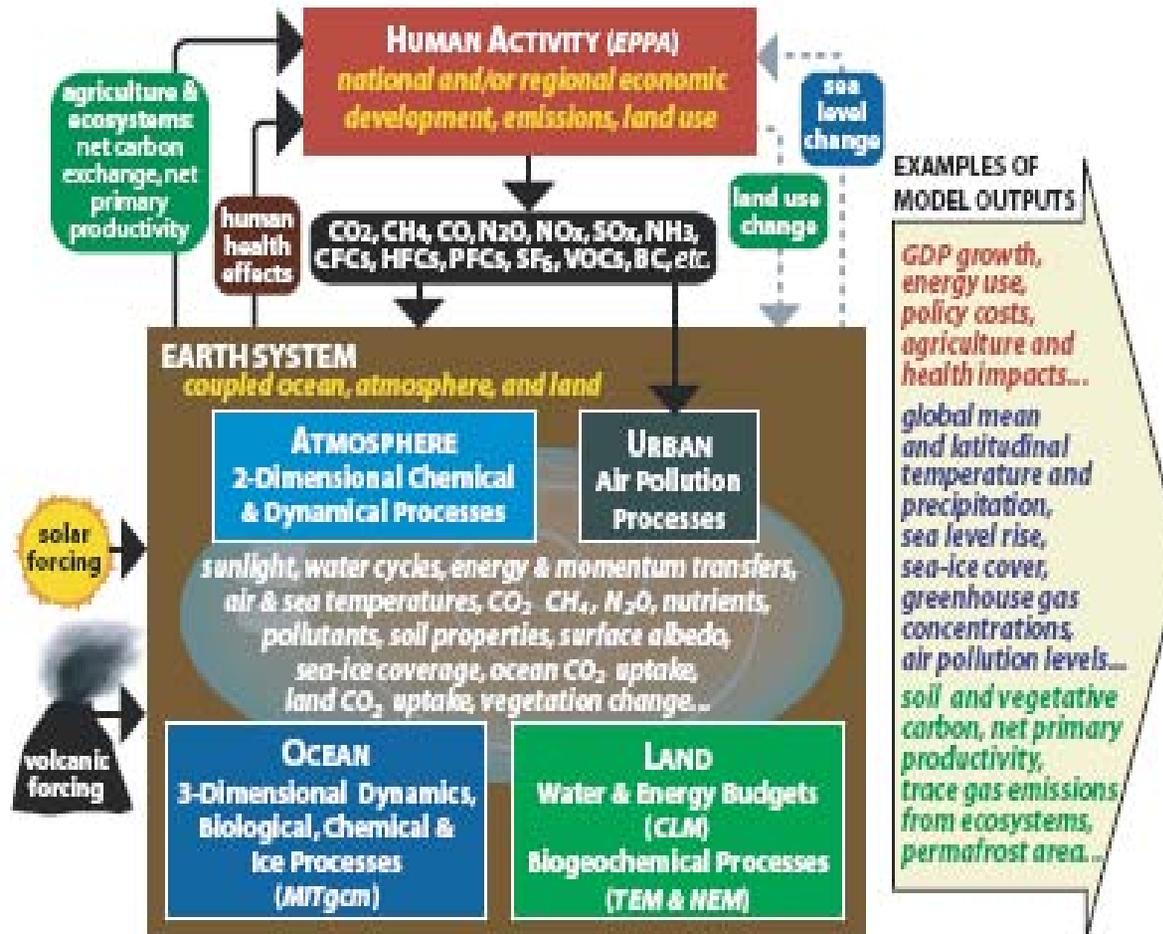
- A model that incorporates the principles of **physics, chemistry, biology** into a mathematical model of climate
e.g. GCM (Global Circulation Model)
- Such a model has to answer what happens to temperature, precipitation, humidity, wind speed and direction, clouds, ice and other variables all around the globe over time



Courtesy of the Intergovernmental Panel on Climate Change. Used with permission.



Example of a climate model : MIT-IGSM



A schematic figure of the MIT-IGSM Version 2

Sokolov *et al.*, 2005



2. DESIGNING a MODEL



Spatial grid

We divide the earth's atmosphere into a finite number of boxes (grid cells).

Assume that each variable has the same value throughout the box.

Write a **budget** for each each box, defining the changes within the box, and the flows between the boxes.

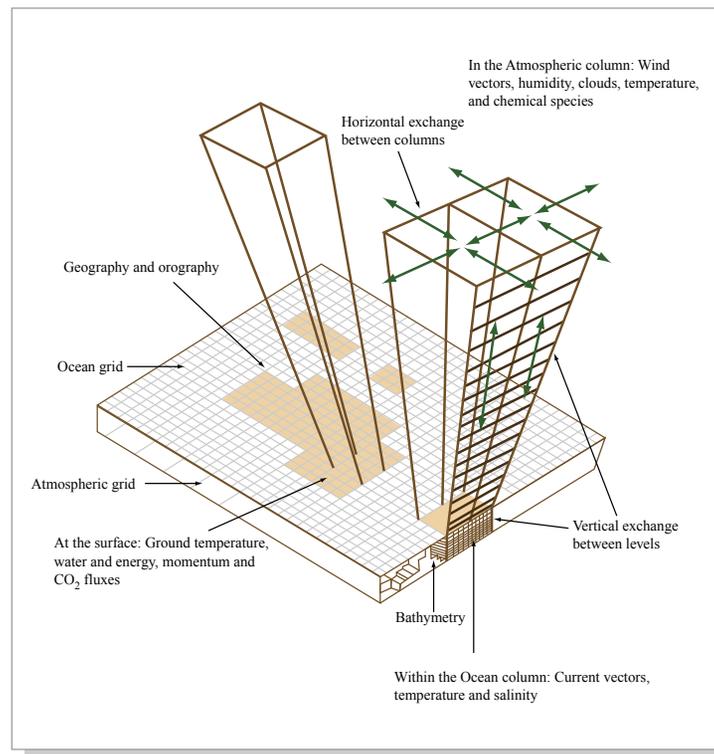


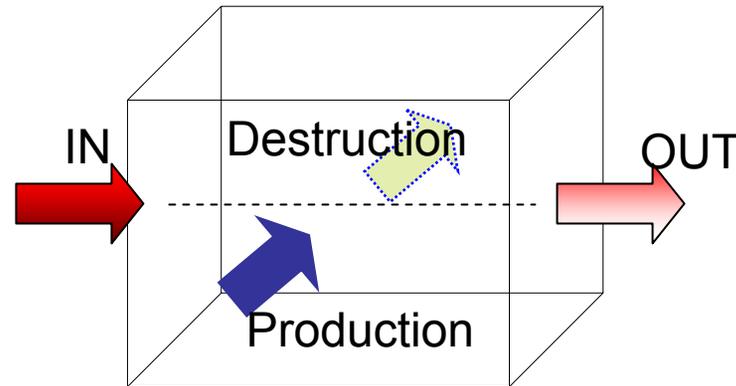
Figure by MIT OpenCourseWare, adapted from Henderson-Sellers: *A Climate Modeling Primer*.



Continuity Equations



We can express changes in a grid cell at a given time step



CHANGE = (Production - Destruction) +/- (Gain or Loss by advection)

$$\left(\frac{\partial \Phi}{\partial t} \right) = \frac{d\Phi}{dt} - \frac{\partial}{\partial x} (u\Phi) - \frac{\partial}{\partial y} (v\Phi) - \frac{\partial}{\partial z} (w\Phi)$$

The total change (rate of accumulation) of Φ in the box

Actual production or destruction of Φ within the box

Change in Φ due to loss to downstream boxes or arrival of Φ from an upstream box (called advection or convection)



Continuity Equations (Atmosphere/Oceans in 3-D)



Solving the Basic Equations for the Atmosphere in 3-D

Mass Continuity

Equations of Motion (momentum continuity)

Thermodynamic Equation (energy continuity)

Chemical Continuity Equation

$$\frac{\partial \rho}{\partial t} = -\frac{d(u\rho)}{dx} - \frac{d(v\rho)}{dy} - \frac{d(w\rho)}{dz}$$

$$\frac{\partial u}{\partial t} = -\frac{d(uu)}{dx} - \frac{d(vu)}{dy} - \frac{d(wu)}{dz} + \dots$$

$$\frac{\partial v}{\partial t} = -\frac{d(uv)}{dx} - \frac{d(vv)}{dy} - \frac{d(wv)}{dz} + \dots$$

$$\frac{\partial w}{\partial t} = -\frac{d(uw)}{dx} - \frac{d(vw)}{dy} - \frac{d(ww)}{dz} + \dots$$

$$\frac{\partial T}{\partial t} = -\frac{d(uT)}{dx} - \frac{d(vT)}{dy} - \frac{d(wT)}{dz} + \frac{1}{c_v} \left(J - p \frac{D(1/\rho)}{Dt} \right)$$

$$\frac{\partial \chi}{\partial t} = -\frac{d(u\chi)}{dx} - \frac{d(v\chi)}{dy} - \frac{d(w\chi)}{dz} + \text{Chemical Production} - \text{Chemical Loss}$$

J : radiation, conduction, latent heat release, etc

$D(1/\rho) / Dt$: conversion between thermal and mechanical energy in fluid system

$\left. \begin{array}{l} + \text{pressure gradient} \\ + \text{Coriolis force} \\ + \text{gravity} \\ + \text{friction} \end{array} \right\}$



The Equations for Chemistry and Biology



- For ocean chemistry, air-sea CO₂ flux is:

$$\begin{aligned} &\text{Air-sea CO}_2 \text{ exchange flux} \\ &= k_s (p\text{CO}_{2,\text{ocean}} - p\text{CO}_{2,\text{atmosphere}}) \end{aligned}$$

- For land biology and chemistry,
atmosphere-land CO₂ flux is:

$$\begin{aligned} &\text{Atmosphere-land CO}_2 \text{ exchange flux} \\ &= \text{Photosynthesis} - \text{Respiration} - \text{Decomposition} \\ &= (\text{GPP} - R_A) - R_H \end{aligned}$$



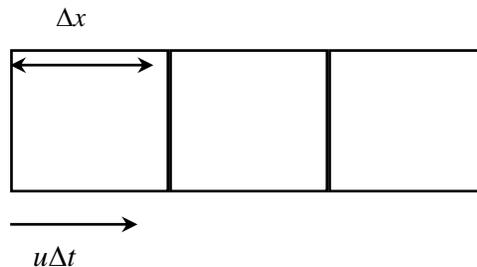
Time stepping and stability



- Time is also treated in discrete units.
- Time intervals depend on the size of the boxes:

General Rule for stability: the CFL condition

$$\frac{u\Delta t}{\Delta x} \leq 1$$



Intuitively don't want to transport more than a grid cell over a time step.

Eg. In atmosphere max $u = 100\text{m/s}$; grid spacing = 300 km;
Constraint: $\Delta t < 3000$ seconds (less than 1 hour)



3. How to Solve the equations



- We want to solve for the values of the variables described by these equations over time.
i.e. to integrate the set of differential equations
- Essentially we have seven (or more) variables described by the same number of equations that describe change with respect to time. (T, p, ρ, u, v, w , water, etc.). So we should be able to solve for the values of the variables through time...
- BUT these equations **cannot be solved analytically**; there is no closed form solution
- So need to use numerics: discretize in time and space...



Demand on Computation



Total Computation Time:

For example, for a 2.8° x 2.8° degree atmospheric model

<u>How Many Grid Cells?</u>	<u>What Happens at each Grid Cell?</u>	<u>How Many Time Steps Per Year?</u>
128 Longitudes	10 Variables	24+ Time Steps per Day
64 Latitudes	* 100 Computations Each	365 Days per Year
* 18 Vertical Levels		
<hr/>	<hr/>	<hr/>
~ 150,000 Grid Cells	~ 1,000 Computations per Grid Cell per Time Step	~ 10,000 Time Steps per Year

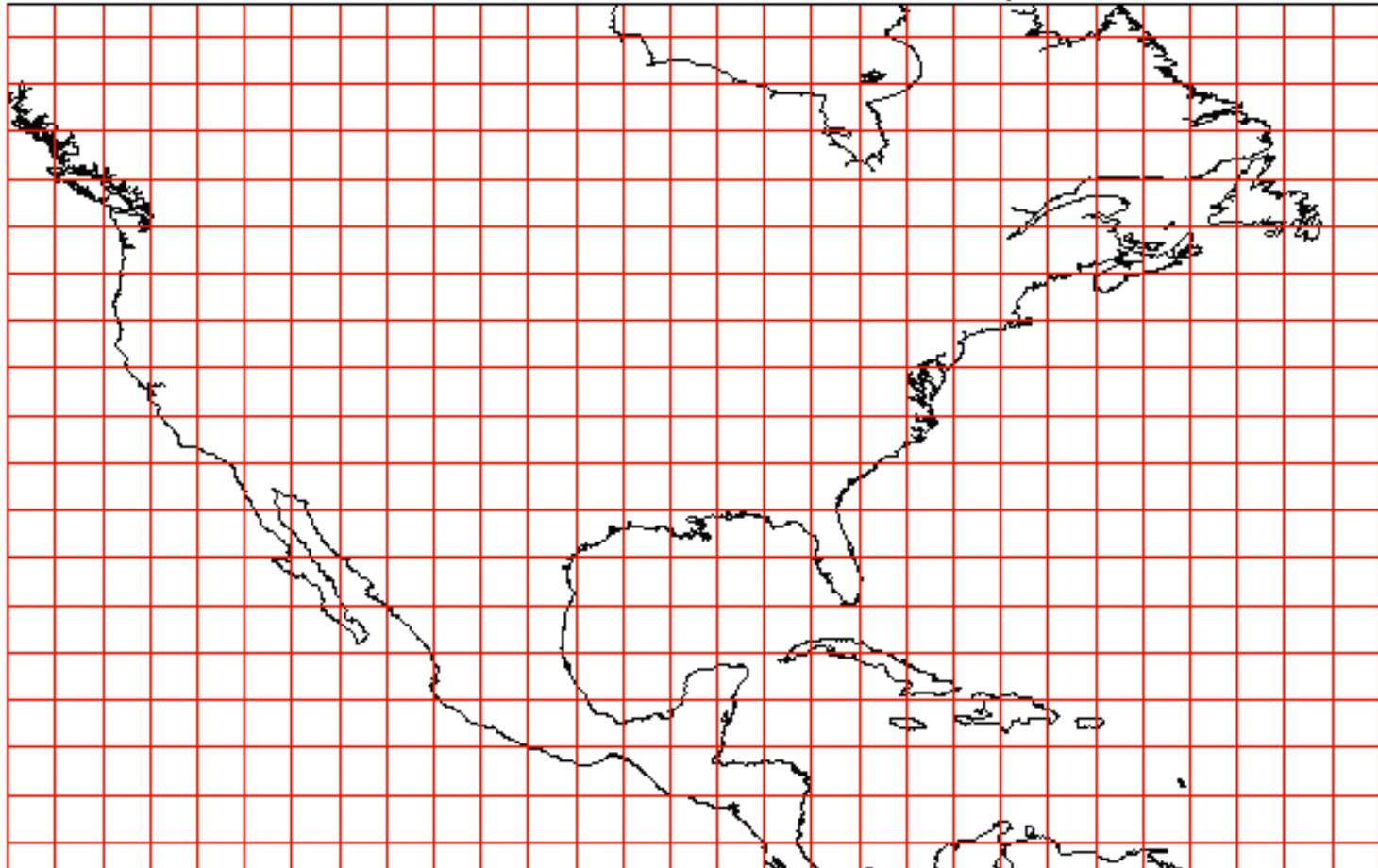
$$150,000 \text{ (Grid Cells)} * 1,000 \frac{\text{Computations}}{\text{(Grid Cell) (Time Step)}} * 10,000 \frac{\text{Time Steps}}{\text{Year}} \approx 1.5 \text{ Trillion} \frac{\text{Calculations}}{\text{Year}}$$

With a 1 GHz machine, a 1 year simulation takes about three hours

And, remember, this is just about the simplest possible model and we generally want to run the model for decades or centuries...



CCM3 Horizontal Resolution (2.8 x 2.8 degrees)





Parameterizations



For GCMs, grid cells are typically hundreds of miles across and often there are thirty vertical layers for the atmosphere.

Many processes happen at smaller scales and must be approximately included (a.k.a., parameterized), including:

- Convection
- Cloud Cover
- Ice Cover: sea and land (glaciers)
- Snow Cover
- Rainfall
- Emissions of Pollutants
- River Runoff into Oceans
- “Eddy Fluxes”
- Sharp weather fronts
- “Gravity Waves”
- Mountains
- Cities (heat islands, emissions, etc)